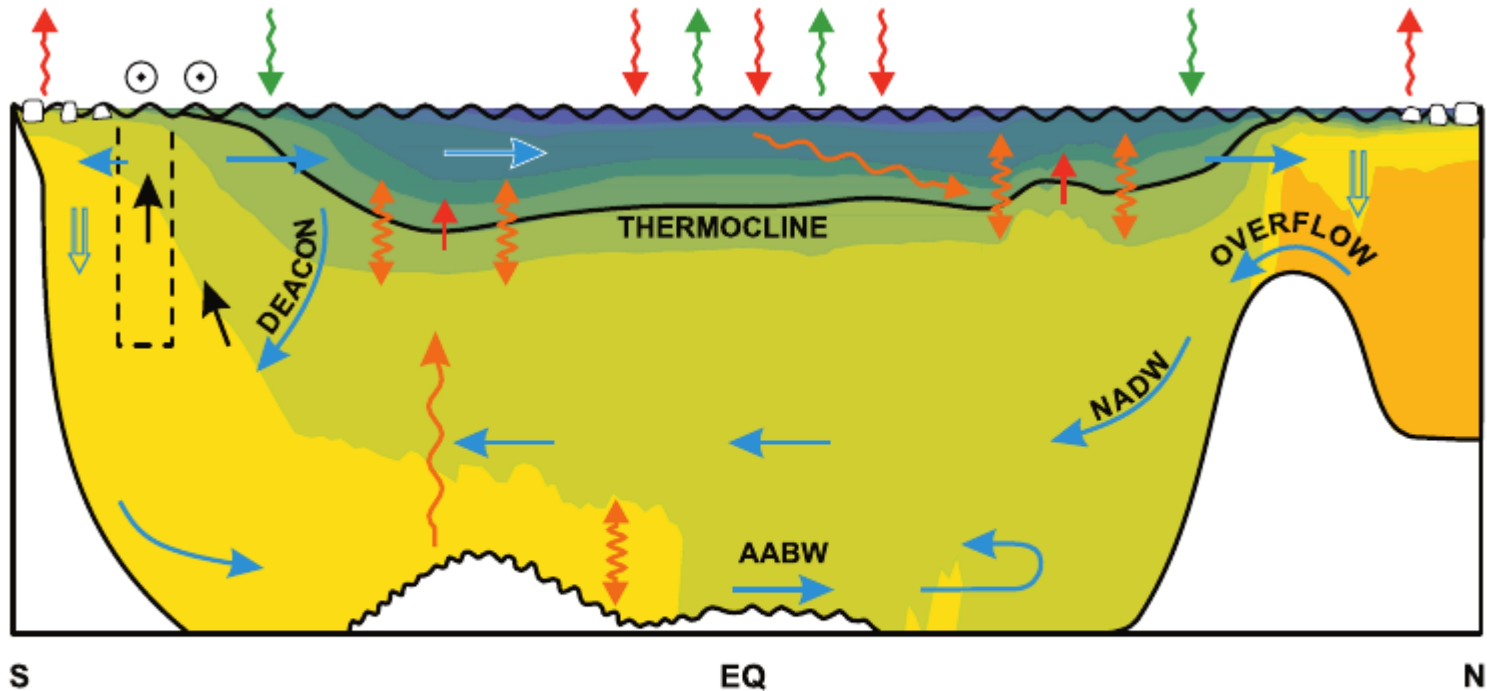













The connection between Southern Ocean Winds, Atlantic Meridional Overturning, and Indo-Pacific Upwelling.

Markus Jochum
Niels Bohr Institute, Copenhagen

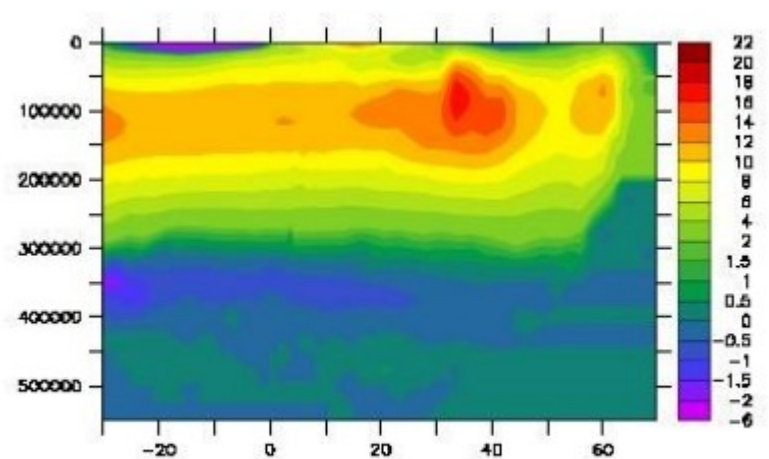
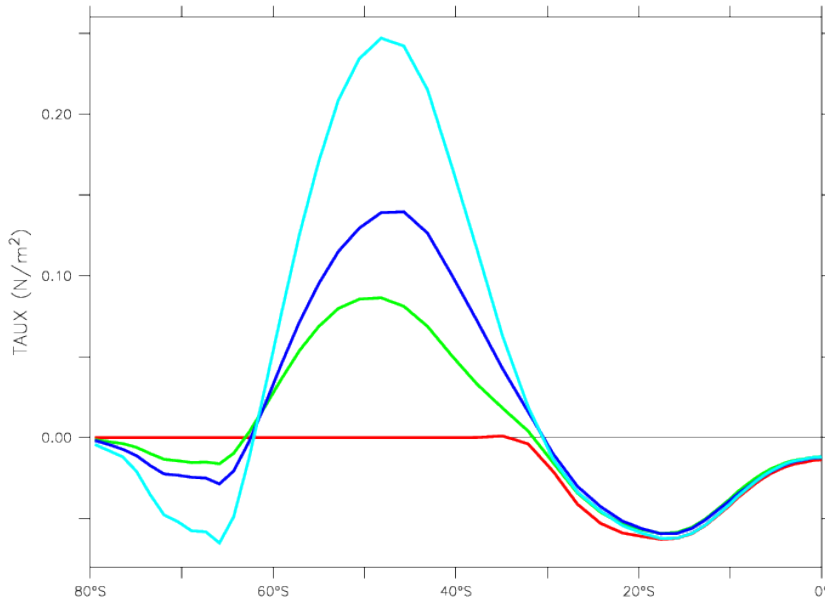
Carsten Eden
Institute for Marine Research,
Hamburg





- | | | | | | |
|---|--------------------------|---|-------------------------|---|----------------------|
|  | volume transport |  | mixing-driven upwelling |  | deep-water formation |
|  | wind-driven upwelling |  | internal waves |  | heat fluxes |
|  | wind |  | diapycnal mixing |  | freshwater fluxes |
|  | profile of Drake passage | | |  | sea ice |

zonally averaged wind stress



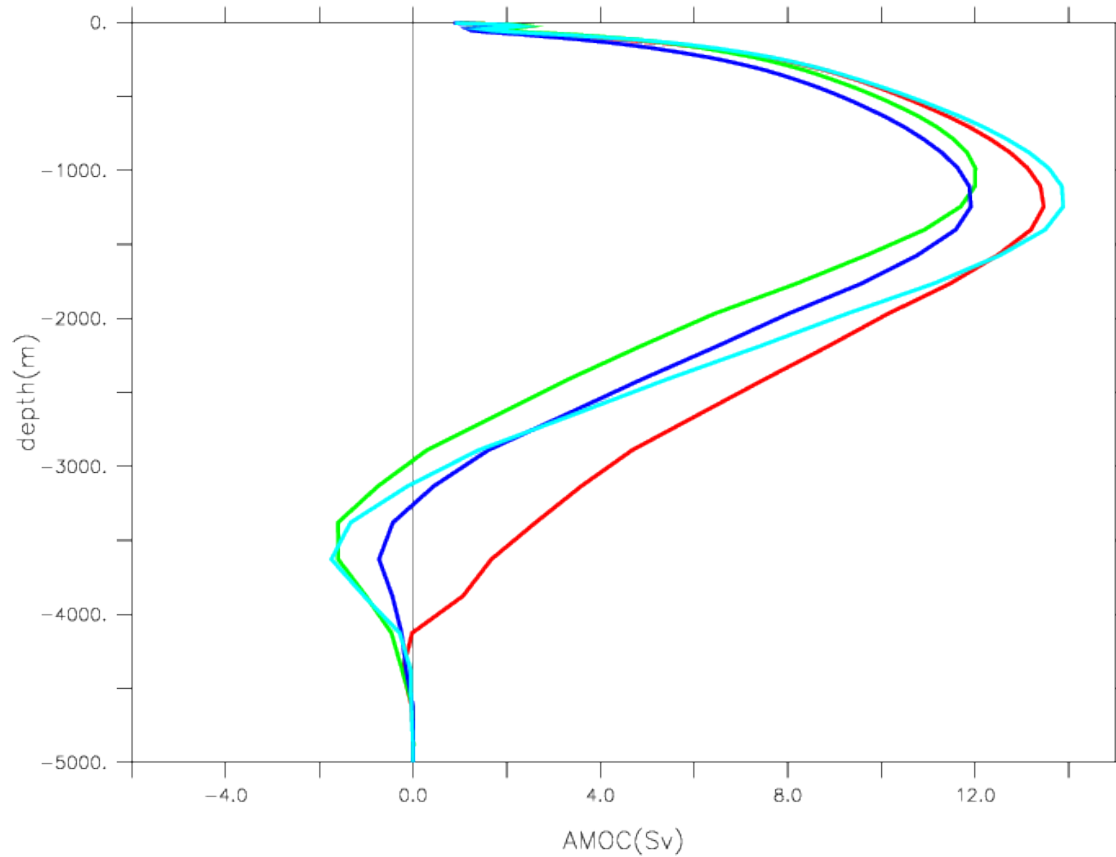
Residual Atlantic Overturning, CONT

Shields et al. 2012

NULL, HALF, CONT, TWO

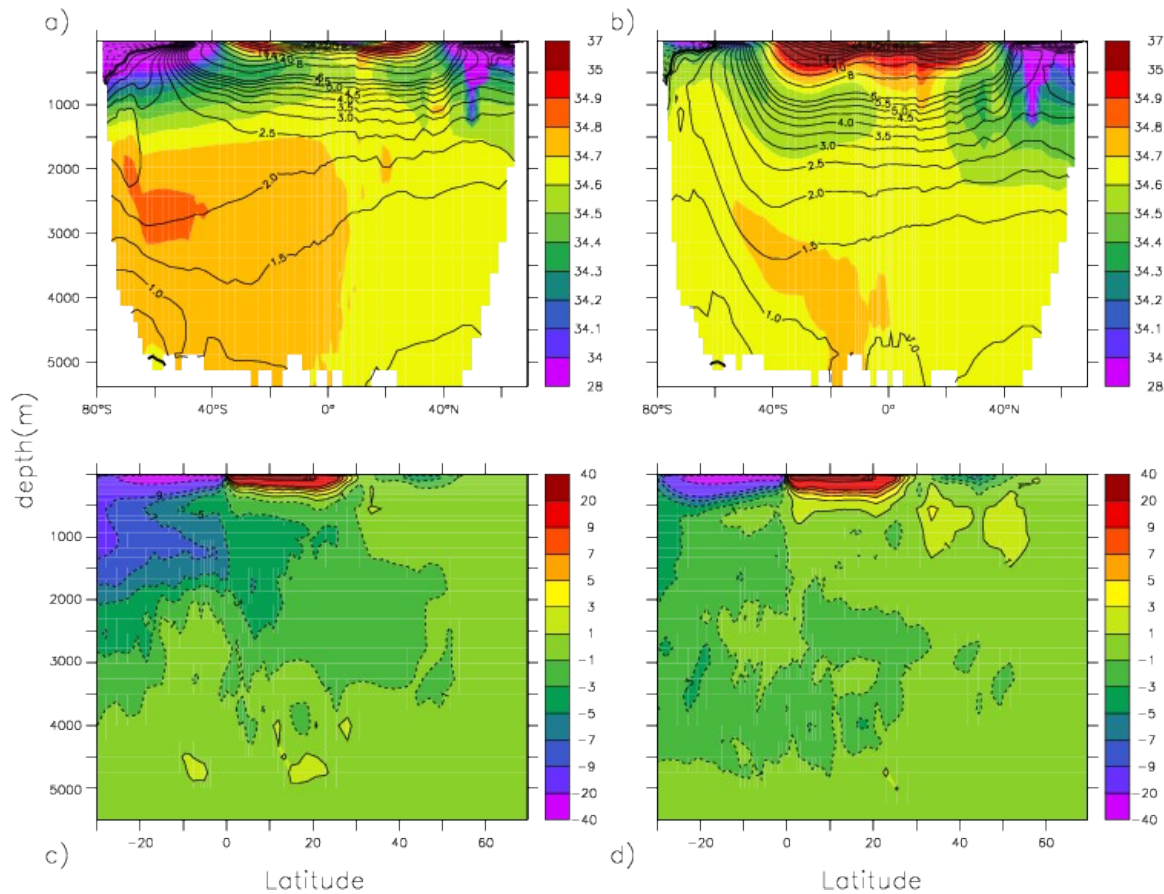
The GCM: CCSM4 in its T31x3 configuration:
 3.75 x 3.75 degree, 26 levels in the atmosphere
 0.6 - 3 degree, 60 levels in the ocean
 Danabasoglu & Marshall the Elder (2007) version
 of stratification dependent isopycnal and thickness diffusion

Residual Overturning at the Equator



NULL, HALF, CONT, TWO

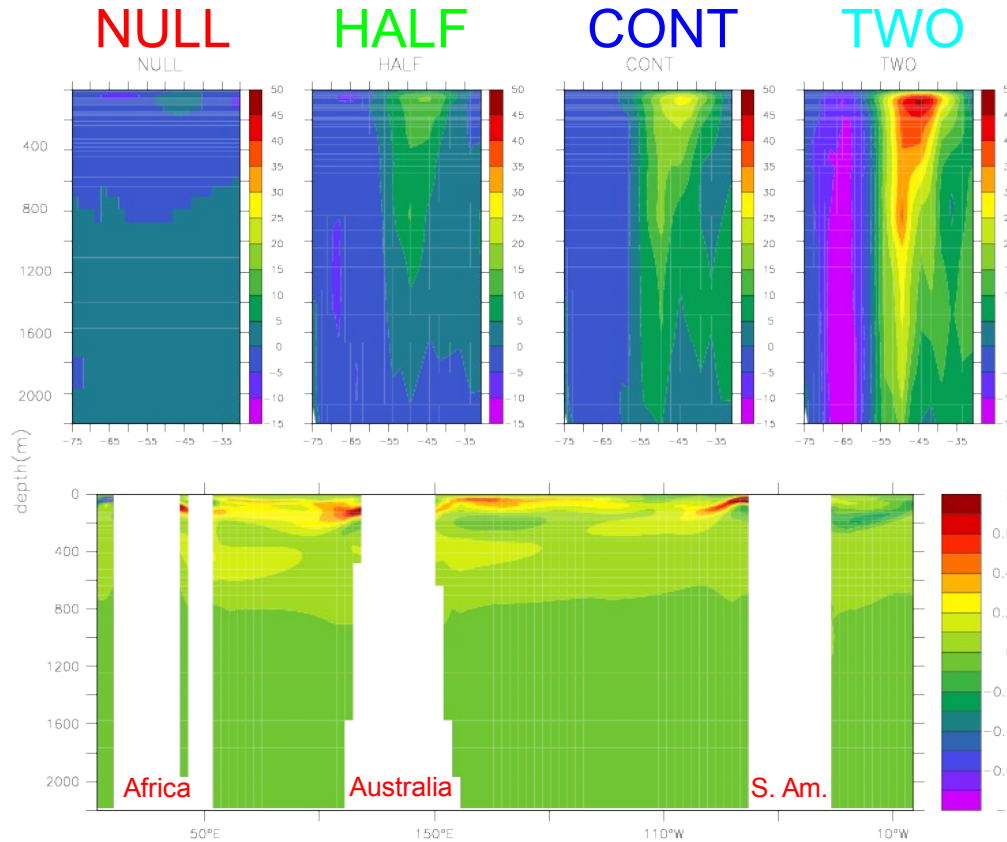
Indo-Pacific Salinity and Temperature



NULL

CONT

Indo-Pacific residual overturning



Southern Ocean
Residual overturning

Stratification along 20S

NULL - TWO

Exp.	τ_{ave}	τ_{DP}	$DP_{tr.}$	$AMOC_{eq}$	Upwelling	Overturning	κ_{GM}
NULL	0	0	0	13	4 / 9	5 / -7 / 4	500
HALF	0.9	0.6	110	12	5 / 7	20 / -11 / 15	570
CONT	1.4	0.6	100	12	9 / 3	35 / -15 / 26	610
TWO	2.5	1.4	180	14	13 / 1	60 / -22 / 42	700

Conclusions

- in CESM there is no eddy compensation or saturation
- in a fully coupled GCM Southern Ocean winds do not affect the Atlantic overturning
- future work to investigate the AMOC should not focus on eddies, but on NADW formation and diapycnal mixing (see Warren 1981 ;-)

