

Application and validation of polynomial chaos methods to quantify uncertainties in simulating the Gulf of Mexico circulation using HYCOM.

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Outline

- Use Polynomial Chaos (PC) expansions to quantify uncertainties in oceanic forecasts
 - **Approximate** model with an accurate surrogate
 - Compute series coefficients via an ensemble
 - Validate the accuracy of the surrogate
 - Mine the series for statistical information in lieu of model
- **Initial Condition uncertainties in the Gulf of Mexico**
 - How to perturb a **field**?
 - How to localize perturbations around a dynamical process?
 - Series Validation
 - Statistical Outputs
- **Uncertainty Experiments to date**
 - Initial conditions uncertainties (small ensemble)
 - Initial conditions and wind forcing uncertainties

HYCOM surrogate idea

- $M(\mathbf{x}, t, \xi)$ is a model output
- ξ is a stochastic variable that represents the dependence of M on the **uncertain** input data
- ξ is characterized by its probability density function $\rho(\xi)$
- The mean of M is: $\bar{M}(\mathbf{x}, t) = \int M(\mathbf{x}, t, \xi) \rho(\xi) d\xi$
- Its variance is

$$\sigma^2(M) = \overline{(M - \bar{M})^2} = \int (M - \bar{M})^2 \rho(\xi) d\xi$$

- A **surrogate** embodies the dependence of M on the uncertain data ξ via a spectral series in ξ :

$$M(\mathbf{x}, t, \xi) = \sum_{n=0}^P \hat{M}_n(\mathbf{x}, t) \psi_n(\xi)$$

$$\text{PC surrogate series } M(\mathbf{x}, t, \boldsymbol{\xi}) = \sum_{n=0}^P \widehat{M}_n(\mathbf{x}, t) \psi_n(\boldsymbol{\xi})$$

- $\widehat{M}_n(\mathbf{x}, t)$: series coefficients
- $\psi_k(\boldsymbol{\xi})$: **orthonormal** basis functions w.r.t. $\rho(\boldsymbol{\xi})$

$$\overline{\psi_m \psi_n} = \int \psi_m(\boldsymbol{\xi}) \psi_n(\boldsymbol{\xi}) \rho(\boldsymbol{\xi}) d\boldsymbol{\xi} = \delta_{m,n}$$

- $\psi_m(\boldsymbol{\xi})$ consists of Legendre polynomials when $\rho(\boldsymbol{\xi})$ is a uniform distribution
- mean: $\overline{M} = \sum_{n=0}^P \widehat{M}_n(\mathbf{x}, t) \overline{\psi_n}, \psi_0 = \widehat{M}_0(\mathbf{x}, t)$
- Variance: $\overline{(M - \overline{M})^2} = \sum_{n=1}^P \widehat{M}_n^2(\mathbf{x}, t)$
- **Where to truncate the series, P ? Monitor Variance**
- **How to determine the coefficients \widehat{M}_n ?**

Calculating the PC series coefficients

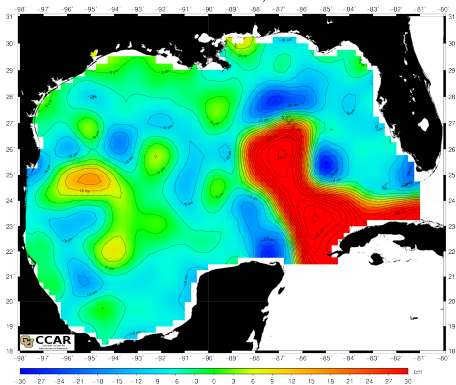
- Minimize the norm of the approximation error by using either Galerkin projection, interpolation, least square, or compressed sensing. All can be implemented via ensemble.
- The least square approaches are useful when model response includes model noise.
- Galerkin projection exploits orthogonality

$$\begin{aligned}\hat{M} = \overline{M\psi_m} &= \int M(\mathbf{x}, t, \boldsymbol{\xi})\psi_m(\boldsymbol{\xi})\rho(\boldsymbol{\xi})d\boldsymbol{\xi} \\ &\approx \sum_{q=1}^Q M(\mathbf{x}, t, \boldsymbol{\xi}_q)\psi_m(\boldsymbol{\xi}_q)\omega_q\end{aligned}$$

- $\boldsymbol{\xi}_q, \omega_q$ are appropriate quadrature roots, weights
- $M(\mathbf{x}, t, \boldsymbol{\xi}_q)$ requires an ensemble at the quadrature roots

Gulf of Mexico Circulation

Historical Mesoscale Altimetry – 05/09/2010



Historical Mesoscale Altimetry – 06/09/2010

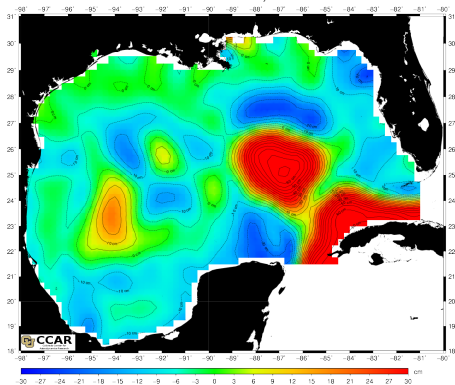


Figure: Sea Surface Height in cm from AVISO

Target Loop Current Eddy separation during May-June 2010.
Capture the role of the frontal anticyclones in eddy detachment.

Uncertainty in Initial Boundary Conditions

- Computational challenge: Number of sample grows exponentially with the number of stochastic variables.
- Rely on EOFs to characterize uncertainty **and** reduce the number of stochastic variables. For 2 EOFs mode we have:

$$M(\mathbf{x}, 0, \xi_1, \xi_2) = \bar{M}(\mathbf{x}, 0) + \left[\sqrt{\lambda_1} \mathcal{M}_1 \xi_1 + \sqrt{\lambda_2} \mathcal{M}_2 \xi_2 \right] \quad (1)$$

- $(\lambda_k, \mathcal{M}_k)$: are eigenvalues/eigenvectors of covariance matrix obtained from free-run simulation
- \bar{M} : unperturbed initial condition
- $M(\mathbf{x}, 0, \xi)$: Stochastic initial condition input
- ξ_1, ξ_2 are the **amplitudes** of the perturbations

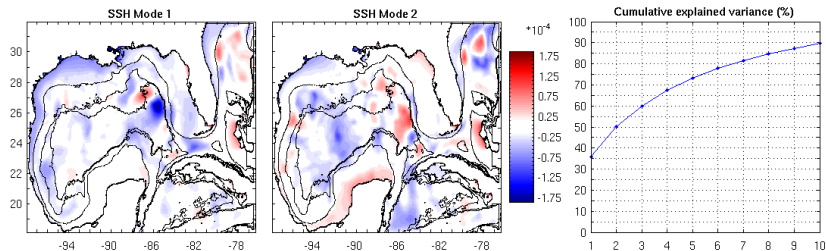


Figure: First and Second SSH modes from a 14-day series. The 2 modes account for 50% of variance during these 14 days.

- Characterize **local** uncertainty: get perturbation from short, 14-day, simulation.
- Uncertainty dominated by Loop Current (LC) dynamics
- Mode 1 seems associated with a frontal eddy

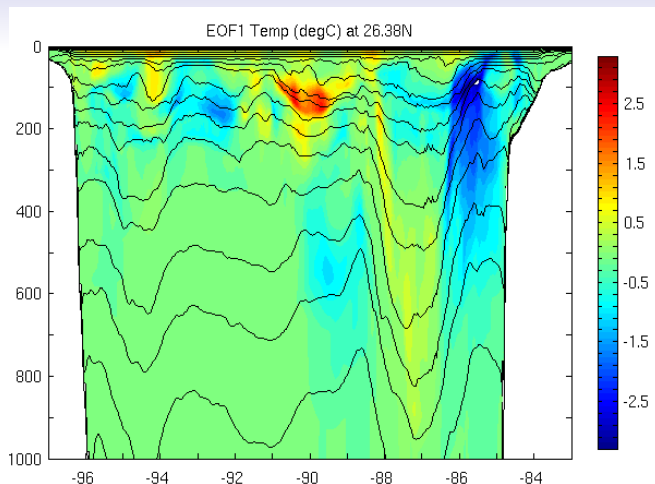


Figure: Vertical slice along 26.4N showing Temperature perturbations. The first mode shows a strong 2.5°C cooling in the vicinity of the frontal cyclones. The "warm" perturbation around 90W is at the southern edge of a small anticyclone NW of the LC.

PC representation

- (ξ_1, ξ_2) independent and uniformly distributed random variables
- PC basis: Legendre polynomials of max degree 6, $P = 28$
- Ensemble of 49 realizations for Gauss-Legendre quadrature

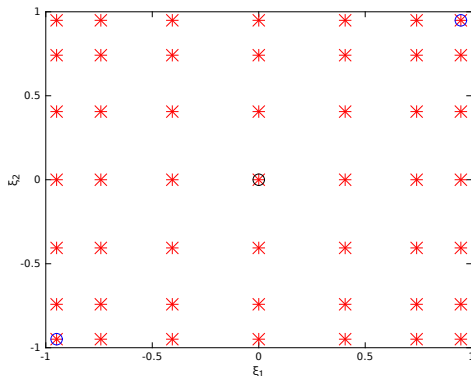
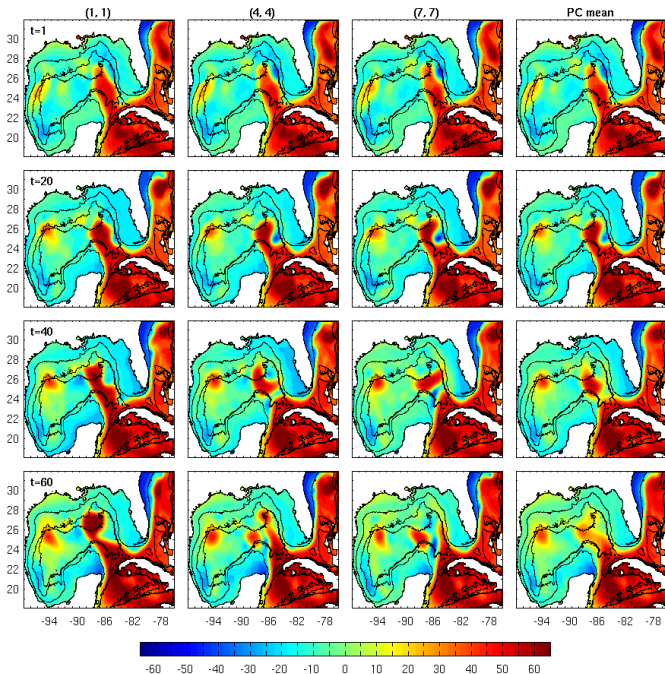


Figure: Quadrature/Sample points in ξ_1, ξ_2 space. Center black circle corresponds to unperturbed run, while blue circles correspond to largest negative and positive perturbations.

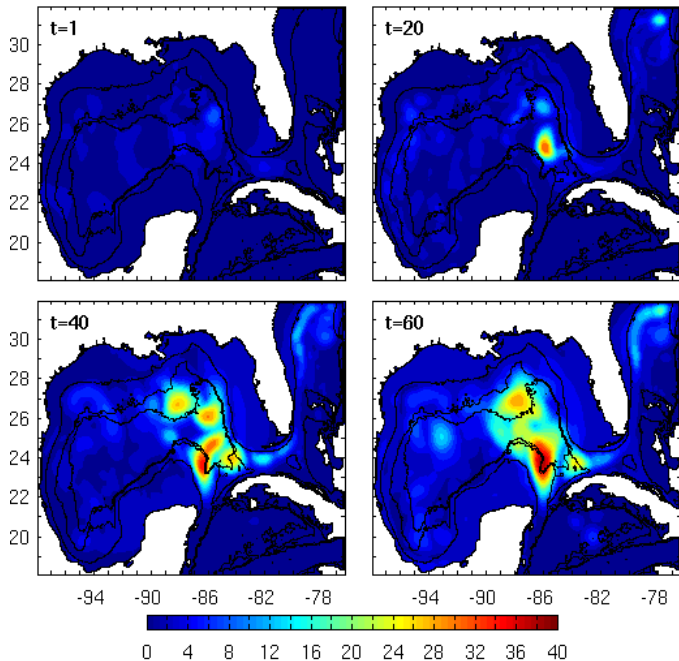


Col 1: SSH of realization (1,1) with weakest frontal eddy

Col 2: SSH of unperturbed realization (4,4) has medium strength frontal eddy

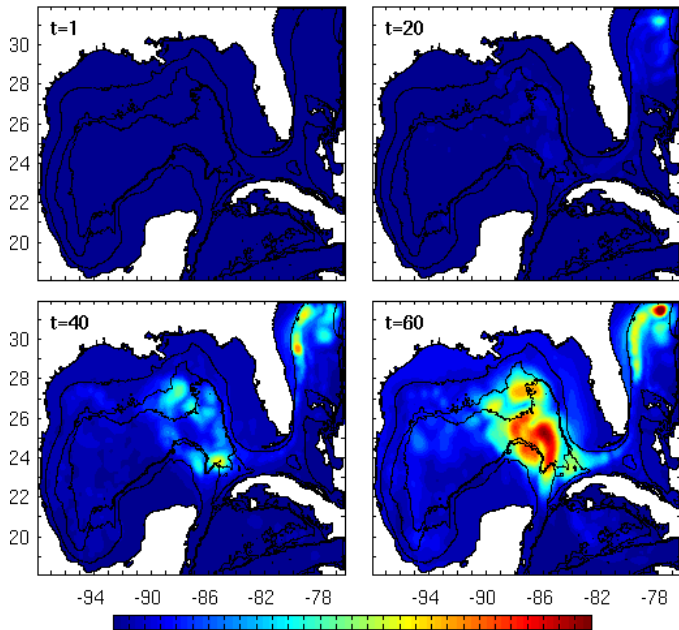
Col 3: SSH of realization (7,7) has strongest frontal eddy and earliest LC separation

Col 4: Loop current edge in ensemble



SSH stddev
(cm) grows in
time with
maximum in
LC region

$$\text{PC-error: } \|\epsilon\|_2^2 = \sum_q [\eta(\bar{x}, t, \xi_q) - \eta_{PC}(\bar{x}, t, \xi_q)]^2 \omega_q$$



SSH
PC-errors
(cm) grow in
time with
maxima in LC
region

On day 60
PC-error is
about 38% of
stddev

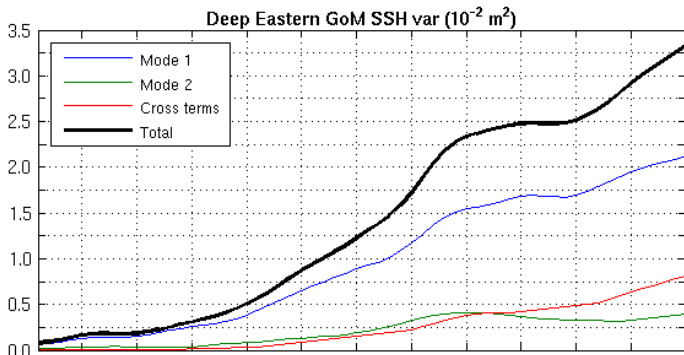


Figure: Variance Analysis: The majority of the variance in the deep part of the Eastern Gulf of Mexico can be attributed to the 1st EOF, while the second mode plays a secondary role, particularly during the time span when the series is reliable (< 40 days, x-axis tick marks interval is 5-days).

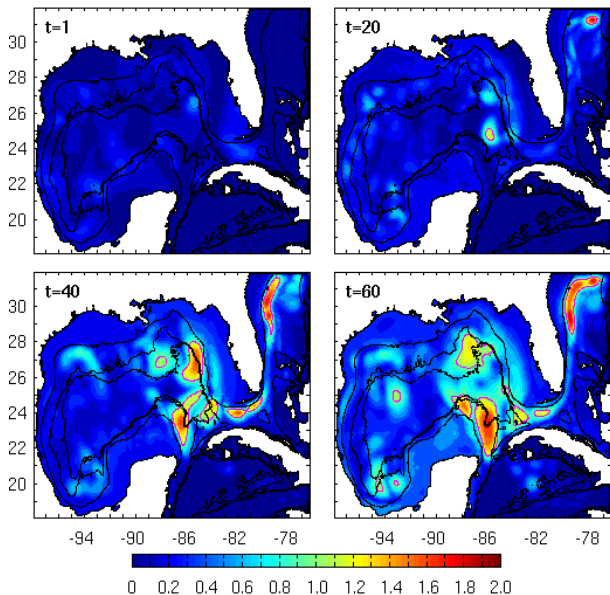


Figure: Predictability Limit: Spatial distribution of the ratio of the forecast standard deviation to climatology standard deviation for SSH (from AVISO). The magenta lines show areas where the ratio > 1 .

Initial Conditions & Wind Forcing Uncertainties

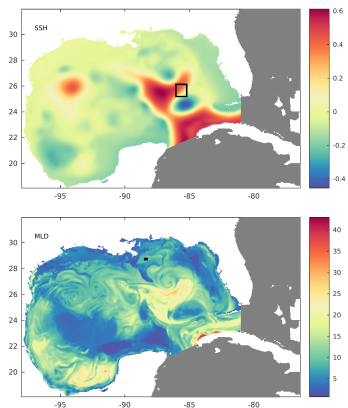


Figure: Field snapshots on day 30: (Top) SSH; (Bottom) MLD

Focus on the following two Quantities of Interest (QoIs):

- Sea Surface Height (SSH) averaged over a square area near the loop current (LC) region:
[-86.04° , -85.20°] in longitude and [25.19° , 26.23°] in latitude
- Mixed Layer Depth (MLD) averaged over the DeepWater Horizon (DWH) region: [-88.44° , -88.28°] in longitude and [28.68° , 28.79°] in latitude

Our HYCOM simulations start from 05-01-2010 to 05-30-2010. Fig.7 on the left shows HYCOM results on the last day of simulation using unperturbed initial and wind forcing fields.

Variance Analysis

1st Order Sensitivity:

$$S_i = \frac{\sum_{\alpha \in \mathcal{S}_1^i} c_\alpha^2 \langle \Psi_\alpha, \Psi_\alpha \rangle}{\sum_{i=1}^P c_i^2 \langle \Psi_i, \Psi_i \rangle}$$

Total Order Sensitivity:

$$\mathcal{T}_i = \frac{\sum_{\alpha \in \mathcal{S}_T^i} c_\alpha^2 \langle \Psi_\alpha, \Psi_\alpha \rangle}{\sum_{i=1}^P c_i^2 \langle \Psi_i, \Psi_i \rangle}$$

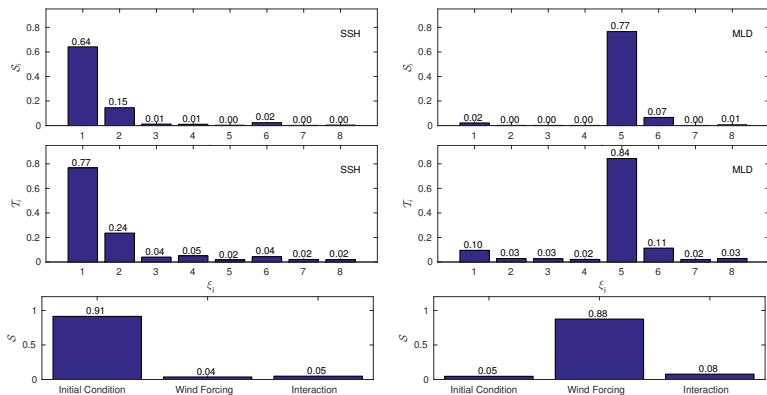
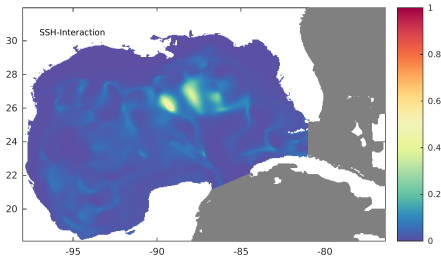
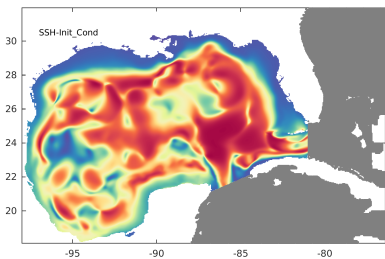


Figure: Sensitivities: (Top) 1st Order; (Middle) total Order; (Bottom) Initial and Wind forcing sensitivities.

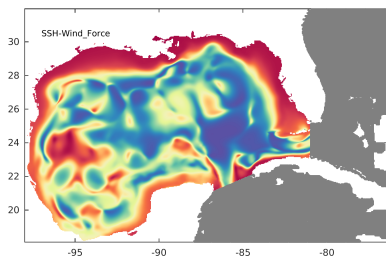
SSH: Joint Sensitivity



(a) Interaction

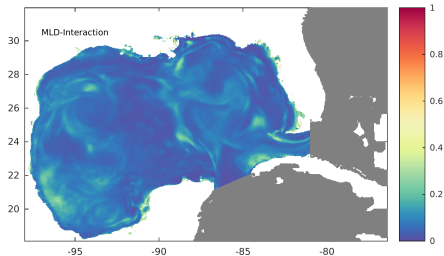


(b) Initial Condition

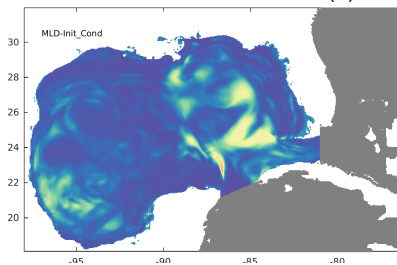


(c) Wind Forcing

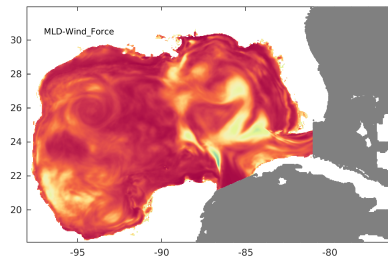
MLD: Joint Sensitivity



(d) Interaction



(e) Initial Condition



(f) Wind Forcing

Conclusions

- Polynomial Chaos are a promising approach in analyzing oceanic uncertainties
- It can monitor representation fidelity and adequacy of sampling
- The 14-day time series EOF analysis helped in designing perturbations focused on a Loop Current separation event.
- The PC-Series performs reasonably well in the 20–40 days range
- Predictability in SSH lost after about 20 days
- It is probably enough to perturb only the 2 leading EOF modes
- Including more uncertainty sources is more important for the current experiment than including more modes.
- SSH uncertainty due mainly to initial conditions except in shelf areas
- MLD uncertainty mainly caused by winds

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- I. Sraj, M. Iskandarani, A. Srinivasan, W. C. Thacker and O. M. Knio, **Drag Parameter Estimation using Gradients and Hessian from a Polynomial Chaos Model Surrogate** *Monthly Weather Review*, **142**, No 2, pp 933941, 2013.