



A joint special edition on

Sources and Sinks of Ocean Mesoscale Eddy Energy

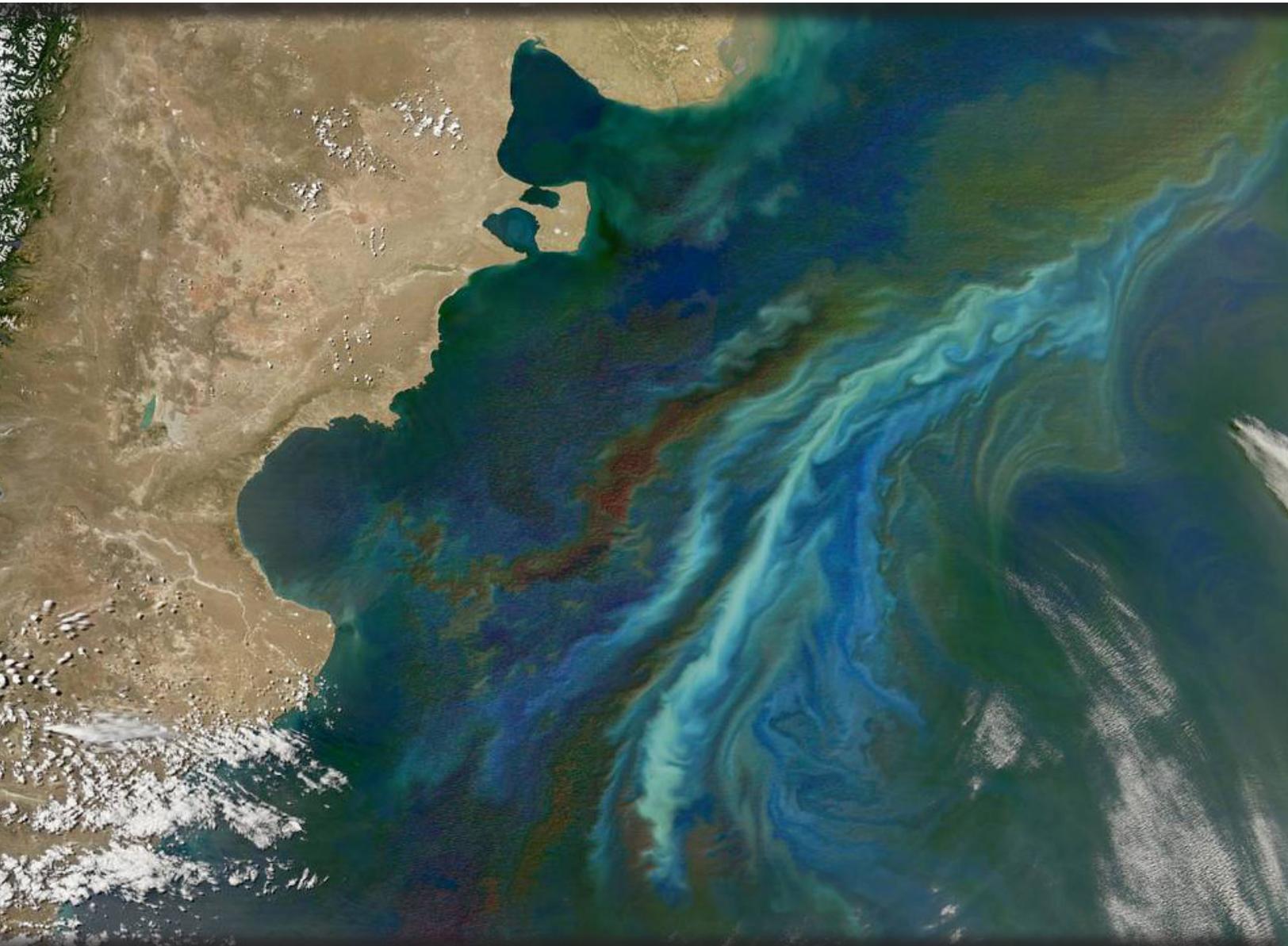


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Impact of horizontal resolution on the energetics of global ocean–sea-ice model simulations

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In order to investigate the origins and consequences of systematic ocean-sea-ice model biases in climate models, an international group of ocean modelers proposed an Ocean Model Intercomparison Project (OMIP; Griffies et al. 2016). The essential element behind the OMIP is a common set of atmospheric and river runoff datasets for computing surface boundary fluxes to drive the ocean-sea-ice models, many of which are used as components of coupled climate system models. The OMIP protocol is an outcome of the Coordinated Ocean-ice Reference Experiments (CORE), which assessed the performance of ocean-sea-ice models (Griffies et al. 2009; Danabasoglu et al. 2014; Griffies et al. 2014; Downes et al. 2015; Farneti et al. 2015; Danabasoglu et al. 2016; Wang et al. 2016a, 2016b; Ilicak et al. 2016; Tseng et al. 2016; Rahaman et al. 2020) using the atmospheric and river runoff dataset of Large and Yeager (2009). However, this dataset has not been updated since 2009 and a new dataset (JRA55-

do; Tsujino et al. 2018) has been developed for the OMIP based on the Japanese Reanalysis (JRA-55) product from Kobayashi et al. (2015) to ensure that it is regularly updated. Tsujino et al. (2020) compares CORE-forced (i.e., OMIP-1) and JRA55-do-forced (i.e., OMIP-2) simulations considering metrics commonly used in the evaluation of global ocean-sea-ice models to assess model biases. Many features are very similar between OMIP-1 and OMIP-2 simulations, but Tsujino et al. (2020) identify many improvements in the simulated fields in transitioning from OMIP-1 to OMIP-2. They attribute many of the remaining model biases either to errors in representing important processes in ocean-sea-ice models (some of which are expected to be mitigated by taking finer horizontal and/or vertical resolutions) or to shared biases in the atmospheric forcing. A first attempt at quantifying the impacts of the models' horizontal resolution on biases was made by Chassignet et al. (2020). They assess

the robustness of climate-relevant improvements in ocean simulations (mean and variability) associated with moving from coarse ($\sim 1^\circ$) to eddy-resolving ($\sim 0.1^\circ$) horizontal resolutions using the same atmospheric forcing dataset (JRA55-do) for both low- and high-resolution configurations. Within the ocean modeling community, it is usually assumed that high-resolution simulations should in general produce better results than low-resolution ones (Fox-Kemper et al. 2019). While this is clearly the case for surface currents and internal variability, greatly enhanced horizontal resolution does not necessarily deliver unambiguous bias improvement in temperature and salinity in all regions (Chassignet et al. 2020). Here, we emphasize some of the salient points of Chassignet et al. (2020) most relevant to the main topic of this special joint issue of US CLIVAR Variations and CLIVAR Exchanges., i.e., “Sources and Sinks of Ocean Mesoscale Eddy Energy.”

Because the goal is to identify the robust differences and improvements associated with increased horizontal resolution given the same forcing datasets, the participating modeling groups configured their high-resolution configuration with similar parameters to that of the coarse-resolution configuration (see Chassignet et al. 2020 for details). The four models that participated in the comparison are the HYbrid Coordinate Ocean Model (HYCOM, Bleck 2002; Chassignet et al. 2003), the ocean (POP) and sea-ice components of the Community Earth System Model version 2 (CESM2, Danabasoglu et al. 2020), the ocean-sea ice component (FESOM) of the

coupled Alfred Wegener Institute Climate Model (AWI-CM, Sidorenko et al. 2015, 2018; Rackow et al. 2018, 2019; Sein et al. 2018), and the LASG/IAP Climate system Ocean Model (LICOM, Zhang et al. 1989; Liu et al. 2004; Liu et al. 2012; Yu et al. 2018; Lin et al. 2020). Because of the large computational cost associated with the high-resolution runs (factor of 1000 more expensive), only one JRA55-do cycle (1958-2018) is analyzed (versus six cycles for the coarse-resolution runs of Tsujino et al. 2020). It is important to note that not all models use the same climatology for the initial conditions, nor do they use the same wind stress formulation (absolute versus relative winds). When evaluating the bias of a numerical simulation, it is performed with respect to the climatology used to initialize the run.

Table 1 shows the domain-averaged mean kinetic energy for all experiments averaged over the last 20 years of the 1958 -2018 integrations. Not surprisingly, the total kinetic energy is significantly higher for the high-resolution experiments over the low-resolution experiments. For the high-resolution configurations, HYCOM has the highest kinetic energy, with a globally averaged value of $\sim 35 \cdot 10^{-4} \text{ m}^2/\text{s}^2$ and LICOM has the lowest kinetic energy, with a globally averaged value of $\sim 15 \cdot 10^{-4} \text{ m}^2/\text{s}^2$ (Table 1). The higher kinetic energy in HYCOM can be partially explained by the wind stress formulation, which does not take into account the ocean current velocities (absolute winds) while the other three models do (relative winds). The latter has an eddy killing effect that can reduce the total kinetic energy by as much as 30% (see Renault et al.

Table 1. Global mean kinetic energy per surface area (units are in $10^{-4} \text{ m}^2/\text{s}^2$).

	Low-resolution	High-resolution
HYCOM	11	36
NCAR	6	24
FESOM	9	19
LICOM	4	14

2019 for a review). This is roughly the difference that is seen here between HYCOM and POP (POP with absolute winds in the wind stress has a level of kinetic energy that is close to HYCOM in Maltrud and McClean 2005, see their Figure 1). But even with the highest resolution used here ($\sim 0.1^\circ$), the total kinetic energy remains significantly lower than what can be inferred from observations and higher resolution models (closer to $50 \cdot 10^{-4} \text{ m}^2/\text{s}^2$, i.e., Chassignet and Xu 2017). The increase in total kinetic energy from the low- to the high-resolution configuration is approximately a factor of three to four for all models, except for FESOM (factor two only). This is probably because the high-resolution FESOM has a highly variable grid spacing (Chassignet et al. 2020) and does not resolve the Rossby radius of deformation everywhere.

The impact on the circulation of increasing the horizontal resolution is two-fold. First, the solution becomes more nonlinear and allows for a better representation of western boundary currents. Second, the first Rossby radius of deformation is resolved throughout most of the domain (Hallberg 2013) and eddies are formed through barotropic and baroclinic instabilities, although higher vertical modes including submesoscale eddies are not often resolved (nor are they resolved by altimetry, i.e. AVISO). Overall, the large-scale patterns are well represented globally, and there is a significant improvement in the representation of the western boundary currents as resolution is increased (Gulf Stream and Kuroshio). In the North Atlantic, most coarse resolution models to date have the tendency to exhibit

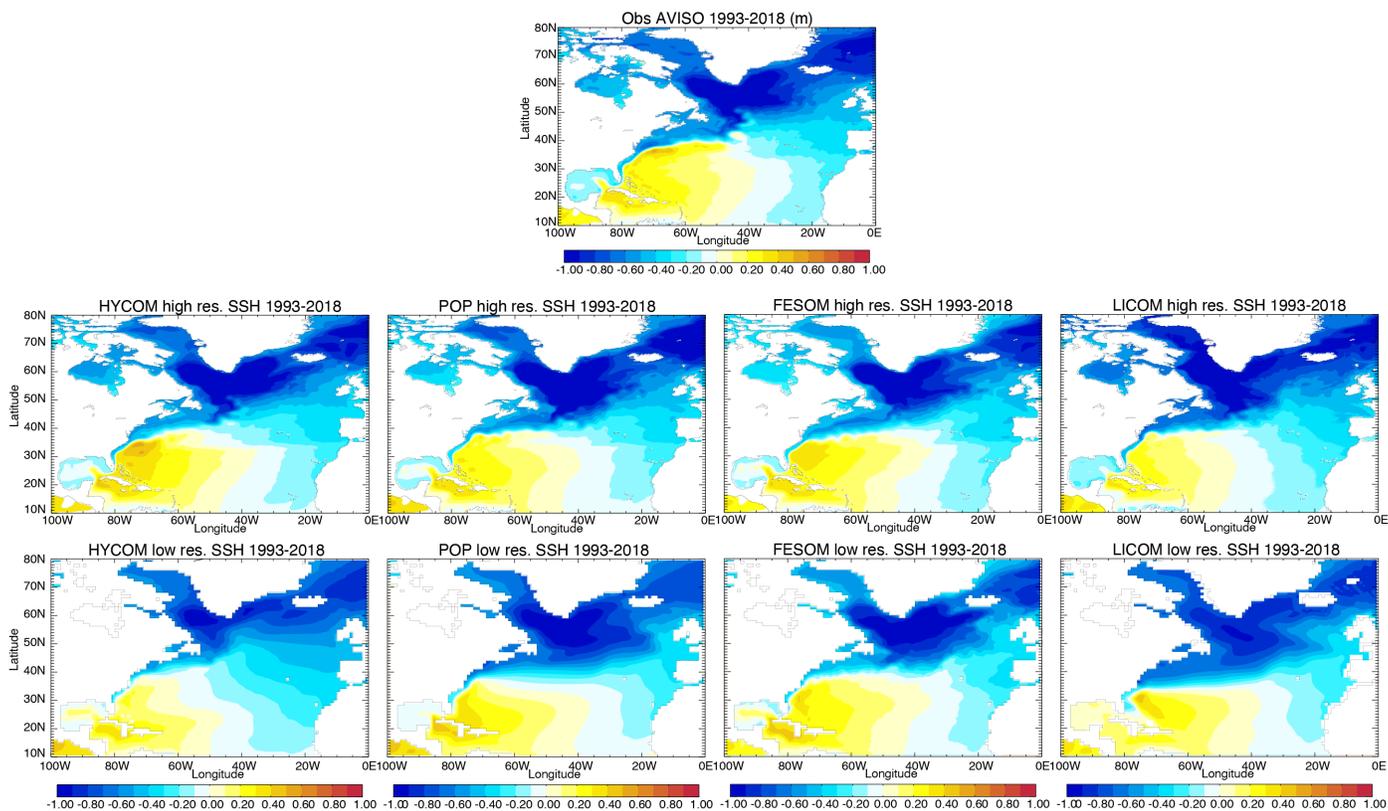


Figure 1. Mean sea surface height fields for observations (Rio et al. 2014) (top panel), high-resolution experiments (middle panel), and low-resolution experiments (lower panel).

an overshooting Gulf Stream and a poor representation of the North Atlantic Current (NAC) at the Northwest Corner (Figure 1). This was the case for three out of the four models. Instead of turning north-northeastward along the continental rise of the Grand Banks and past the Flemish Cap to a latitude of $\sim 51^\circ\text{N}$ before turning eastward (see review paper by Rossby 1996), the modeled NAC is strongly zonal in POP, FESOM, and LICOM, and does not turn northward near the Grand Banks (Figure 1). This has been a long-standing issue for many ocean components of the CMIP climate models, and it does not necessarily improve as the computational mesh is refined. Increasing the horizontal resolution does improve the Gulf Stream separation (see Chassignet and Marshall 2008 and Chassignet and Xu 2017 for a review) in all models, but not necessarily the representation of the Northwest Corner circulation. HYCOM is the only model that had a good representation of the Northwest Corner in both the low- and high-resolution experiments (Figure 1). Since all the models use the same atmospheric forcing dataset, the difference is solely due to the numerical and physical choices made by each modeling group.

As expected, there is a significant increase in the sea surface height (SSH) variance as resolution increased and the eddying solution SSH variance maps are much closer to the observations than their low-resolution counterpart (Chassignet et al. 2020). The surface eddy kinetic energy (EKE) maps for two of the high-resolution simulations (HYCOM and POP) and from observations (AVISO) are displayed in Figure 2. The AVISO EKE map has some inherent smoothing since the along-track measurements were optimally interpolated on a 0.25° grid which filters scales less than 150 km (due to track separation and measurement noise and errors) and time scales less than 10 days (repeat cycle of the altimeters). To match the observations, the modeled EKE maps are computed using 10-day average outputs and this time averaging removed much of the small-scale variability associated with inertial motions and ageostrophic effects (Chassignet and Xu 2017). Overall, the EKE is larger in HYCOM because absolute winds are used to force the models, but in all high-resolution experiments ($\sim 0.1^\circ$), the

variability is still lower than observed, especially in the gyre interiors and in the experiments that used relative winds (POP, FESOM, and LICOM). This underestimation is thus partly a consequence of the eddy-killing effect which results from considering the shear between atmospheric wind and ocean current when computing the wind stress, which can reduce the kinetic energy by as much as 30% (see above discussion). However, the use of absolute wind in HYCOM is not sufficient to raise the level of surface variability to that of the observations, and Chassignet and Xu (2017) argue that one actually needs to significantly increase the resolution ($\sim 0.01^\circ$) in order to resolve the submesoscale instabilities that can energize the mesoscale (Callies et al. 2016) and therefore enhance EKE comparable to the mesoscale AVISO observations. It is more physical to take into account the vertical shear between atmospheric winds and ocean currents when computing the wind stress (see Renault et al. 2019, for a review), as it allows for a better representation of western boundary current systems (Ma et al. 2016), especially the Agulhas Current retroflexion and associated eddies (Renault et al. 2017). In HYCOM, the Agulhas eddies are too regular and follow the same pathway (Figure 2). The use of relative winds does improve the pathway for the Agulhas current eddies (NCAR vs. HYCOM), but it also reduces the level of EKE in the Antarctic Circumpolar Current (ACC) and suppresses variability in many areas. This is especially true for the Indian Ocean, in the tropics, and west of the Hawaiian Islands. Most model-observations comparisons usually focus on the surface fields because of the scarcity of long time series at depth covering a large spatial area. While the EKE at depth in the high-resolution experiments is a significant improvement over the quasi non-existent EKE of the coarse-resolution simulations, it is still significantly less than very limited observations (Richardson, 1993; Ollitrault and Colin de Verdière 2014). There is however much less EKE at depth in the experiments using relative winds (POP, FESOM, or LICOM) than in the experiment using absolute winds (HYCOM, not shown). Significantly higher resolution may be necessary in order to obtain a level of EKE at depth close to the observations (see Chassignet and Xu 2017 for a discussion).

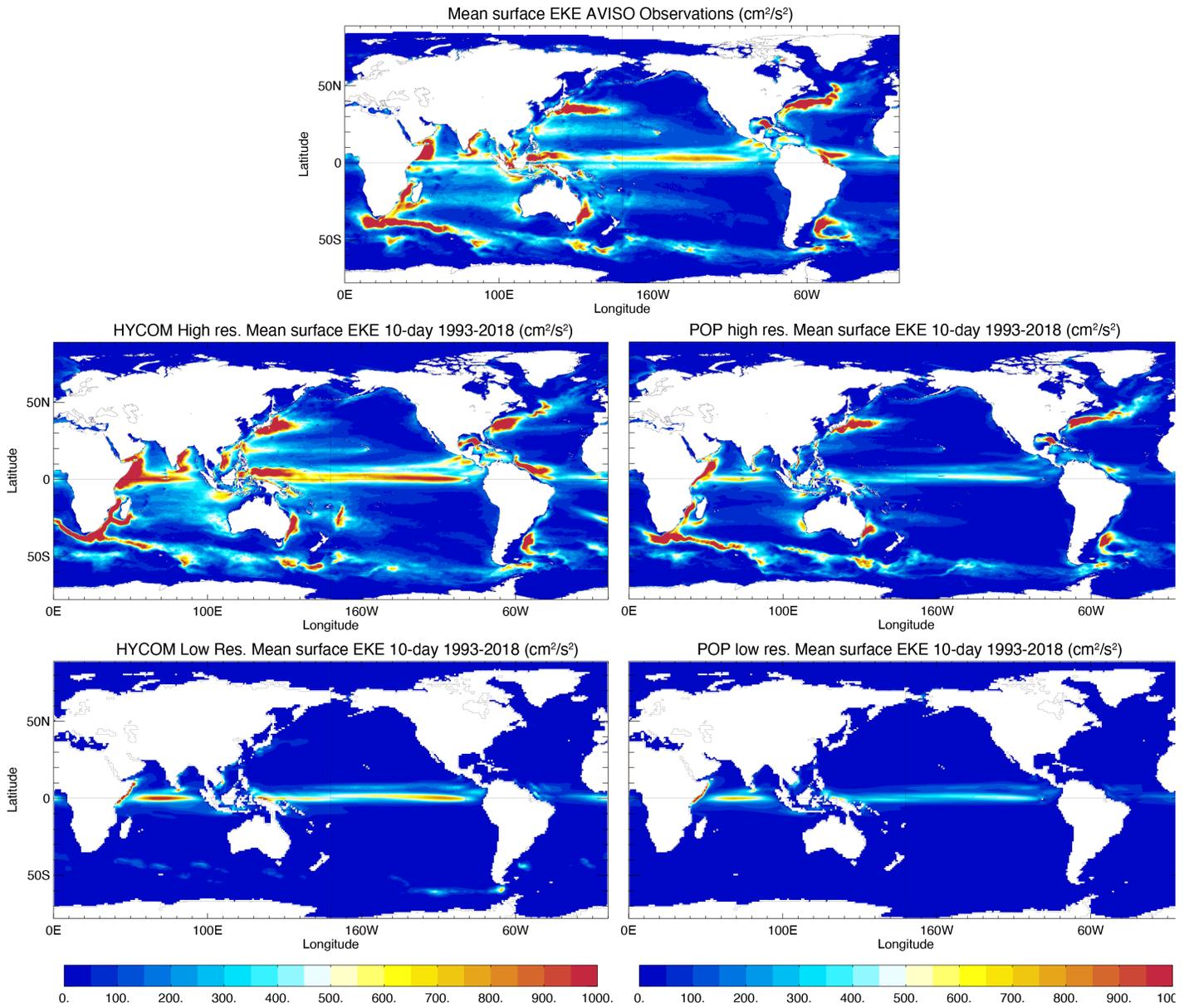


Figure 2. Mean surface eddy kinetic energy for observations (top panel), high-resolution experiments (middle panel), and low-resolution experiments (lower panel).

Overall, the gross features of the bias patterns in low resolution models – position, strength, and variability of western boundary currents, equatorial currents, and ACC – are significantly improved in the high-resolution models. However, despite the fact that the high-resolution models “resolve” these features, the improvements in temperature or salinity are inconsistent among different model families and some regions show increased bias over their low-resolution counterparts (see Chassignet et al. 2020 for a discussion). SSH variability and near-surface EKE are significantly – even qualitatively – improved in all high-resolution models over their low-resolution counterparts, although all of these models still underpredict the observed SSH variability and EKE, particularly in the ocean interior, which indicates a need for further refinements in resolution (Chassignet and Xu 2017) and improvements in less dissipative subgrid schemes for high-resolution models (Pearson et al. 2017). The results in coupled models in the HighResMIP ensemble (Haarsma et al. 2016) show similar improvements in SSH and SST variability and EKE. Considerable differences in the high-resolution models used here are associated with the use of relative winds versus absolute winds.

Another interesting aspect of the high-resolution models versus the low-resolution models is that the interannual variability in the ACC, Indonesian Throughflow, and Atlantic Meridional Overturning Circulation (AMOC) transports (not shown) is more consistent among the high-resolution models than among the low-resolution models (Chassignet et al. 2020). Consistency in all of these transports potentially indicates that higher-resolution models are needed to represent process variability, which may explain some of the past difficulties in comparing the magnitude of these phenomena across coarse-resolution models. However, the mean ACC transport and AMOC strength are not in greater agreement among the high-resolution than the low-resolution models, which means that more work remains in evaluating sensitivity to numerics and subgrid-scale schemes for high-resolution models. Furthermore, Danabasoglu et al. (2016) note that low-resolution models come into greater agreement in AMOC variability after more cycles of the

CORE forcing. This comparison is limited by the cost of the high-resolution models to only a single cycling of the forcing. The short duration of a single forcing cycle limits the comparison of the decadal changes that are emphasized in Danabasoglu et al. (2016), so the improved agreement among the high-resolution models is year-by-year rather than decade-by-decade. Nonetheless, the high-resolution models have systematically stronger and more variable AMOC, in better agreement with observations, both in maximum overturning and profile, than the low-resolution models.

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