Simulating Tides in the Gulf of Mexico using the HYbrid Coordinates Ocean Model

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In order to implement tides in a semi-enclosed basin one needs to prescribe:

- in the whole domain, a **Local Tidal Potential** forcing
- at the limits of the domain, an **Open Boundaries** forcing, which represents the tidal flow coming from outside the domain.
The Local Tidal Potential forcing is applied to the whole domain.
Introduction: Open Boundaries (Remote Forcings)

The Open Boundaries forcing is applied locally.
Introduction

Gouillon et al., 2008

- Simulation of tides in the Gulf of Mexico (GOM) assuming a barotropic ocean and using the Navy Coastal Ocean Model
- Tides are best described when Open Boundaries and Local Tidal Potential forcings are combined
- Bottom friction is the only mechanism of energy dissipation taken into account

Goals of the present study

1. to validate the simulation of barotropic tides in the GOM with HYCOM
2. to consider an additional mechanism of energy dissipation: the conversion from barotropic to baroclinic modes
3. to quantify the loss of energy due to the internal wave mixing
run HYCOM with a barotropic ocean configuration and assess its ability to model barotropic tides. Comparison against:
- previous numerical experiment with a barotropic ocean using NCOM
- data assimilative tidal model in the Gulf of Mexico (Kantha’s, 2000)

run HYCOM with a baroclinic ocean (realistic stratification) and evaluate the loss of energy by tidal conversion process. This should be done with or without bottom friction to isolate the conversion process
Methods: HYCOM

- Vertical coordinate system combines: isopycnal, pressure level (z), terrain following (σ) coordinates.
- Discretization along isopycnal surfaces provides physically more realistic flows and prevents spurious diapycnal mixing.
- η is assumed to be null in the pressure gradient \( \Rightarrow \) limitation in the coastal areas.
Methods

Model configuration

- domain: Gulf of Mexico + Atlantic
- resolution: 1/25°, i.e. 541 × 481 grid points
- barotropic configuration ⇒ 3 layers, \( O(\Delta \rho) \approx 10^{-3} kg.m^{-3} \)
- full isopycnal coordinates
- model is run for 15 days (spin-up is reached after 4 days)
- the last 4 days are used to compute tidal amplitudes and phases

Boundaries conditions:

- Tidal forcing at Open Boundaries obtained with an inverse global model (Egbert’s TPX06.2 model)
- Flather condition at Open Boundaries
- Bottom friction
Preliminary results:

**Comparison of phases and amplitudes for $M_2$ and $O_1$ tides between HYCOM, NCOM and observed tides**
In HYCOM the near-coast amplitudes are too weak in the West part of the basin

The position of the amphidromic point is slightly shifted to the east
A different shape of the coamplitude lines 0.05 m in the middle of the basin (Yucatan Channel Mexico coast).

General phase lag in the basin and absence of reflection pattern in the Louisiana bay.
Kantha’s model for $O_1$ tides

HYCOM simulation for $O_1$ tides

- Good agreement of tidal amplitudes in the main part of the domain
The amphidromic point is shifted to the S-E
NCOM has higher tidal amplitudes
Preliminary results: Limits of our barotropic approximation

- No propagation of internal tides because of the barotropic ocean configuration
- A realistic density gradient would generate internal gravity waves on steep topography
Conclusion

So far the simulation is promising. However more exploration on tidal phases and local tidal potential implementation need to be done. Some of the differences could be explained by:

- bathymetry smoothing and resolution (too coarse)
- geometry of the basin